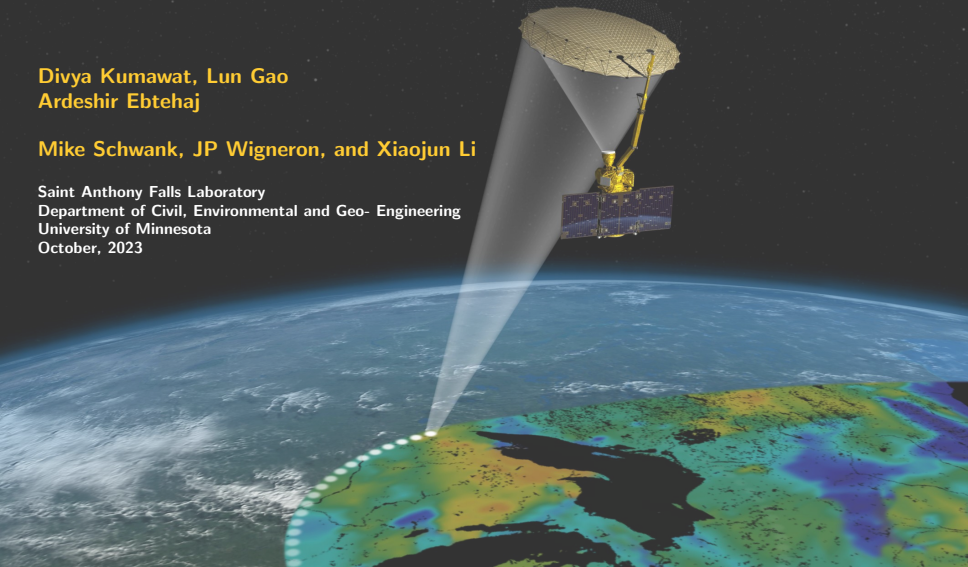


# Land Surface Heterogeneity and High-resolution Retrievals of L-band VOD and Soil Permittivity over the Arctic Boreal Forest and Permafrost

**Divya Kumawat, Lun Gao  
Ardeshir Ebtehaj**

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October, 2023



# Outline

**Introduction**

**Current Research**

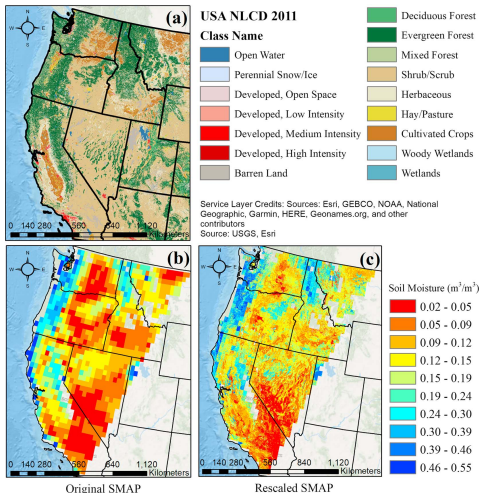
**Future Research**



## High-resolution SMAP Retrievals

- ▶ A physical or a data-driven problem?

$$SM = f(\mathbf{x})$$



## Constrained Multi-Channel Algorithm (CMCA)

- ▶ Accounting for slow changes of VOD in time

$$\psi_t^* = \underset{\psi_t}{\operatorname{argmin}} \sum_{k=1}^K \sum_{t=0}^T \|\mathbf{E}_{kt}^{-1/2} [\mathbf{e}_{kt}^p - f_{\tau-\omega}(\psi_t)]\|_2^2 + S(\psi_t) \text{ subject to } \psi_t^l \preceq \psi_t \preceq \psi_t^u$$

$$S(\psi_t) = \lambda_0 \|\theta_t^p\|_2^2 + \sum_{i=1}^n \lambda_i \|\mathbf{D}_i \tau_t\|_2^2$$

- ▶ Porosity at 1 km (Miller and White, 1998)



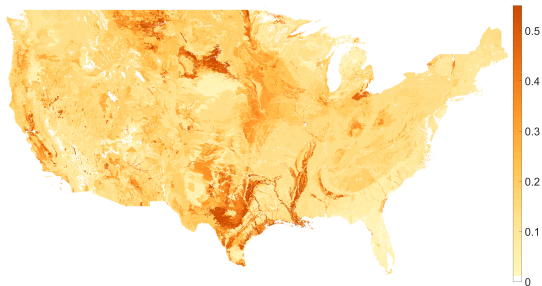
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- ▶ Clay fraction at 1 km (Miller and White, 1998)



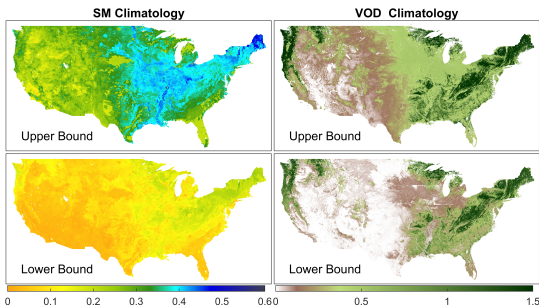
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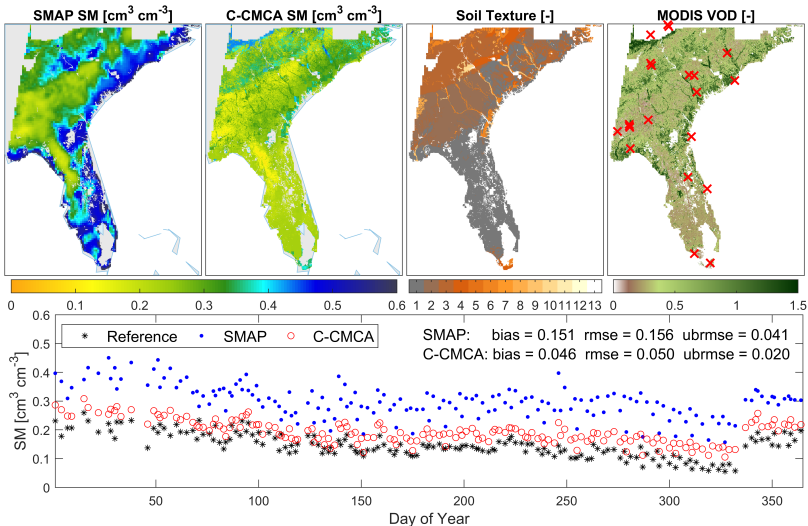
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- ▶ NLDAS 2006–2015 (12.5 km) and MODIS VOD data from 2000–2016 (1 km)



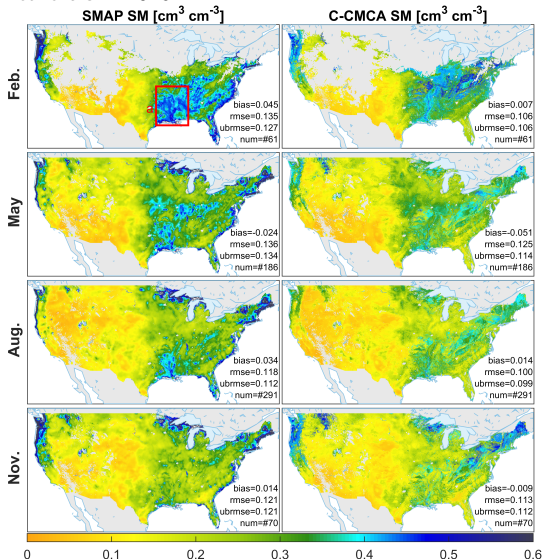
# SMAP Retrievals

## ► SM snapshot retrievals – Feb 15, 2016



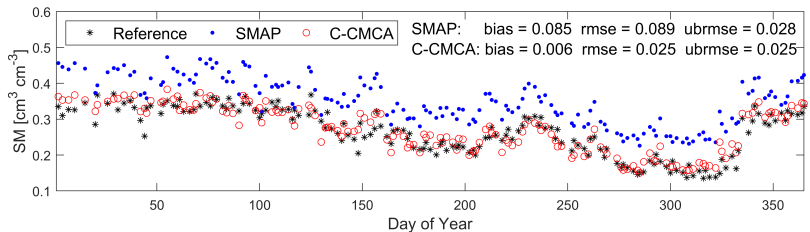
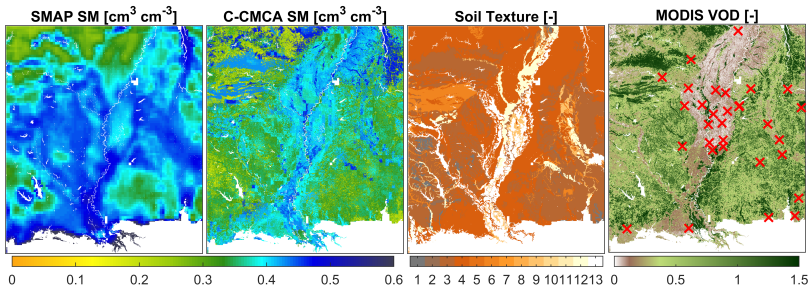
## SMAP Retrievals

► SM monthly retrievals in 2016

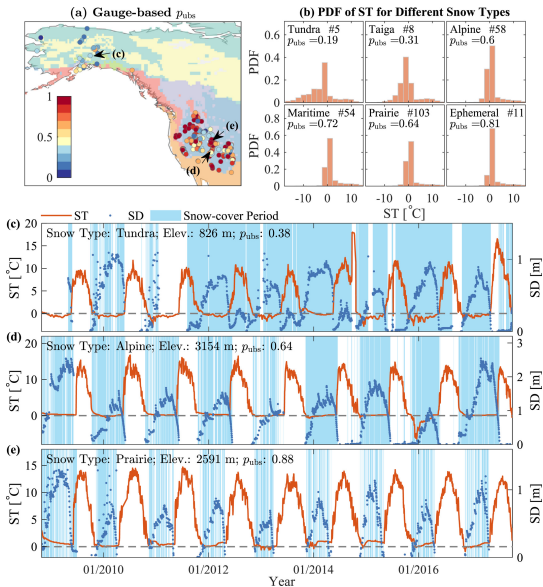


# SMAP Retrievals

## ► SM monthly retrievals in February 2016



# Soil Moisture below Snowpack





## Soil-Snow-Canopy Radiative Transfer Model in L-band

$$Tb^P = \overbrace{T_g e^P \gamma_v}^{(1)} + \overbrace{T_c (1 - \omega) (1 - \gamma_v)}^{(2)} + \overbrace{T_c (1 - \omega) (1 - \gamma_v) r^P \gamma}^{(3)}$$

$e^P$ : effective emissivity of soil-snow system

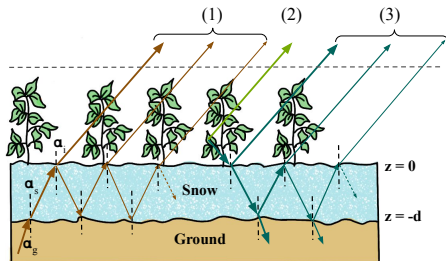
$r^P$ : effective reflectivity of soil-snow system

$T_g$ : ground temperature

$T_c$ : canopy temperature.

$\gamma_v$ : vegetation transmissivity.

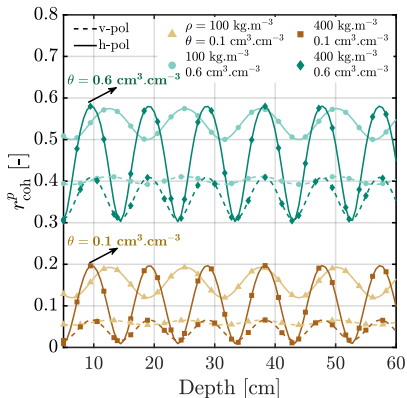
$\omega$ : vegetation single scattering albedo.



Kumawat D., M. Olyaei, L. Gao, A. Ebtehaj, Passive Microwave Retrieval of Soil Moisture below Snowpack at L-band using SMAP Observations, IEEE Trans. on Geosci. and Remote Sens., DOI:10.1109/TGRS.2022.3216324.

## Soil-Snow-Canopy Radiative Transfer Model in L-band

$$r_{\text{coh}}^p = \frac{\xi_{\text{CS}}^p + \tilde{\xi}_{\text{SG}}^p e^{-2\gamma_s d_s \cos \alpha_s}}{1 + \xi_{\text{CS}}^p \tilde{\xi}_{\text{SG}}^p e^{-2\gamma_s d_s \cos \alpha_s}} \quad \text{and} \quad e_{\text{coh}}^p = 1 - \frac{\tilde{\xi}_{\text{SG}}^p + \xi_{\text{CS}}^p e^{-2\gamma_s d_s \cos \alpha_s}}{1 + \xi_{\text{CS}}^p \tilde{\xi}_{\text{SG}}^p e^{-2\gamma_s d_s \cos \alpha_s}},$$



$$r_{\text{inc}}^p = 1 + \text{sgn}(|\xi_{\text{CS}}^p| |\tilde{\xi}_{\text{SG}}^p| - 1) \frac{|\xi_{\text{CS}}^p|^2 + |\tilde{\xi}_{\text{SG}}^p|^2 - |\xi_{\text{CS}}^p|^2 |\tilde{\xi}_{\text{SG}}^p|^2 - 1}{|\xi_{\text{CS}}^p|^2 |\tilde{\xi}_{\text{SG}}^p|^2 - 1}$$

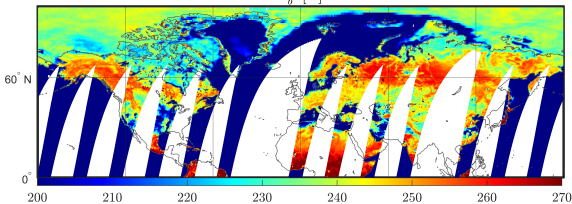
# Retrieval of High-latitude VOD and Ground Permittivity

- ▶ SMAP Orbit: 23 Jan 2017

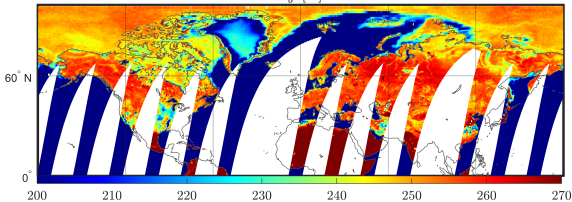
$$\phi^* = \underset{\phi}{\operatorname{argmin}} \sum_p \left( y_{TB}^p - f^p(\phi) \right)^2 + \mu (\tau - \tau_0)^2$$

subject to  $\phi_l \leq \phi \leq \phi_u$ ,

$T_b^h$  [K]



$T_b^v$  [K]

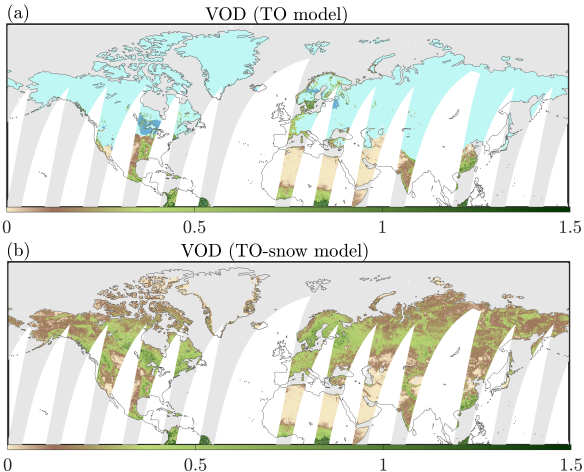


# Retrieval of High-latitude VOD and Ground Permittivity

► VOD retrievals

$$\phi^* = \underset{\phi}{\operatorname{argmin}} \sum_p \left( y_{TB}^p - f^p(\phi) \right)^2 + \mu (\tau - \tau_0)^2$$

subject to  $\phi_l \leq \phi \leq \phi_u$ ,



# Retrieval of High-latitude VOD and Ground Permittivity

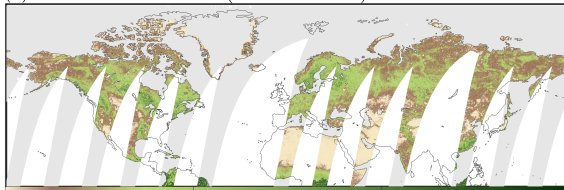
► Soil permittivity retrievals

$$\phi^* = \underset{\phi}{\operatorname{argmin}} \sum_p \left( y_{TB}^p - f^p(\phi) \right)^2 + \mu (\tau - \tau_0)^2$$

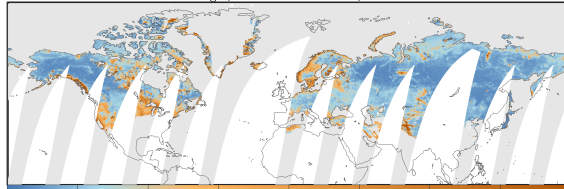
subject to  $\phi_l \leq \phi \leq \phi_u$ ,

VOD (TO-snow model)

(b)



0 0.5 1 1.5  
(c)  $\epsilon_g$  (TO-snow model)



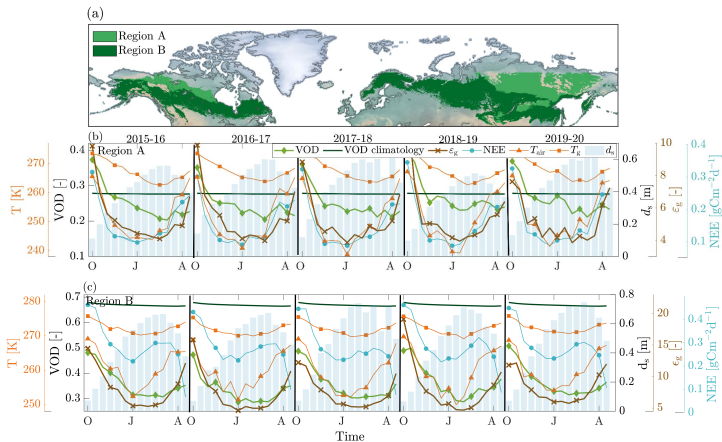
0 5 10 15 20 25 30 35 40

# Retrieval of High-latitude VOD and Ground Permittivity

- Boreal forests with continuous (Region A) and sporadic permafrost (Region B)

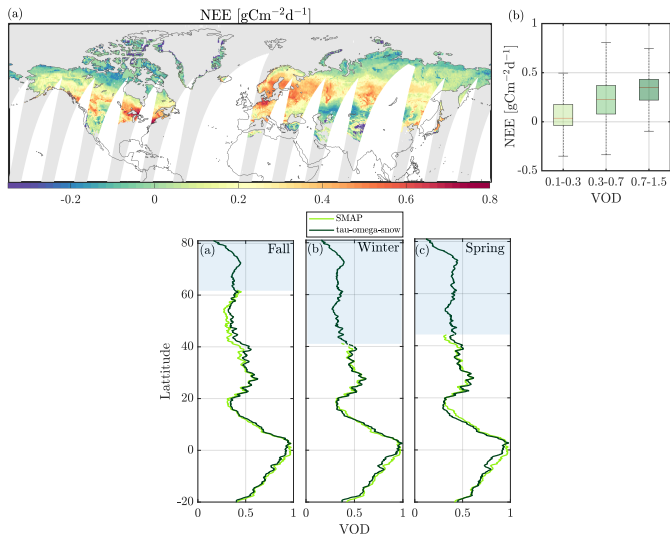
$$\phi^* = \underset{\phi}{\operatorname{argmin}} \sum_p \left( y_{T_B}^p - f^p(\phi) \right)^2 + \mu(\tau - \tau_0)^2$$

$$\text{subject to } \phi_l \leq \phi \leq \phi_u,$$



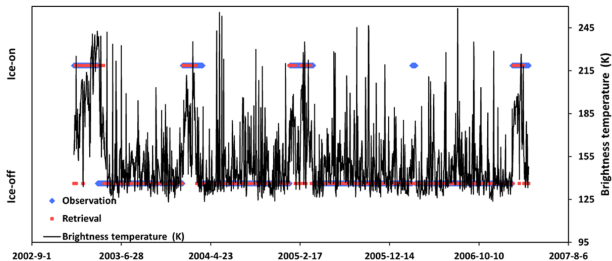
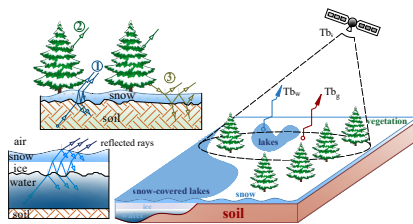
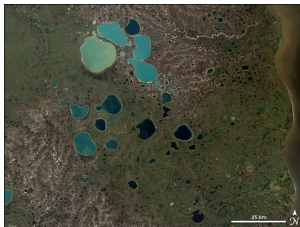
## Retrieval of High-latitude VOD and Ground Permittivity

- FLUXCOM NEE in 01/23/17 and VOD.



# Arctic Freeze-Thaw Dynamics

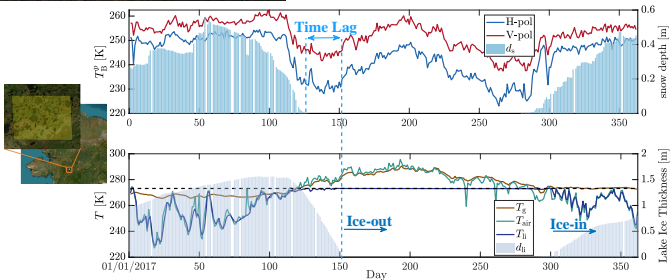
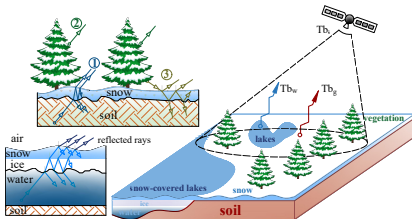
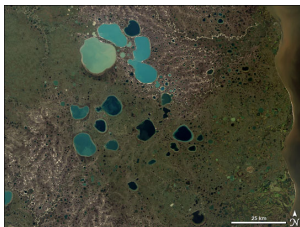
$$Tb_w^p = T_w e_{is}^p \Rightarrow Tb^p = f_w Tb_w^p + (1 - f_w) Tb_g^p$$



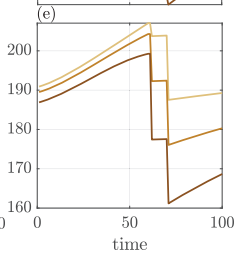
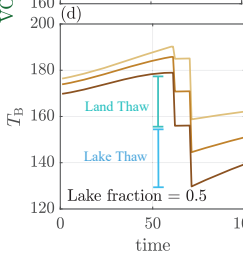
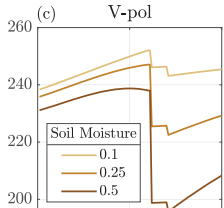
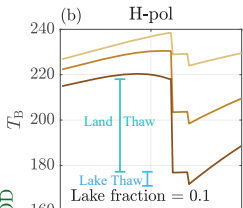
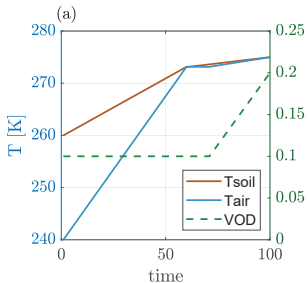


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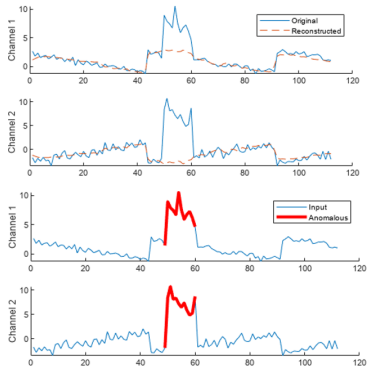
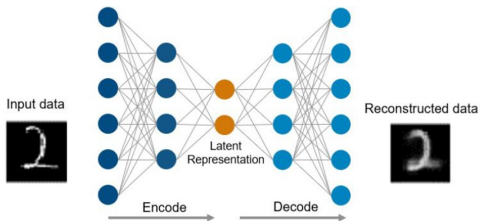
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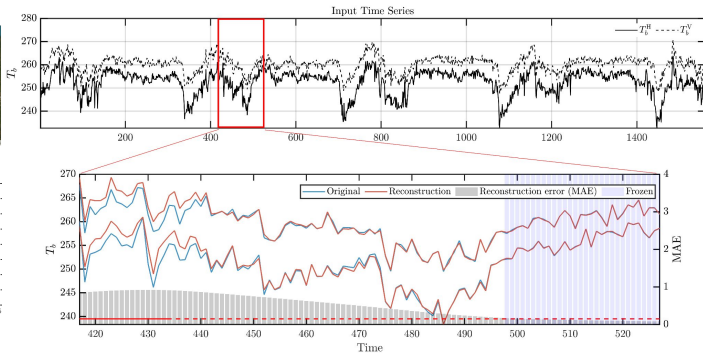
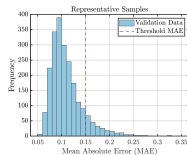
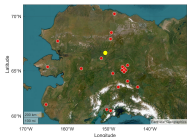
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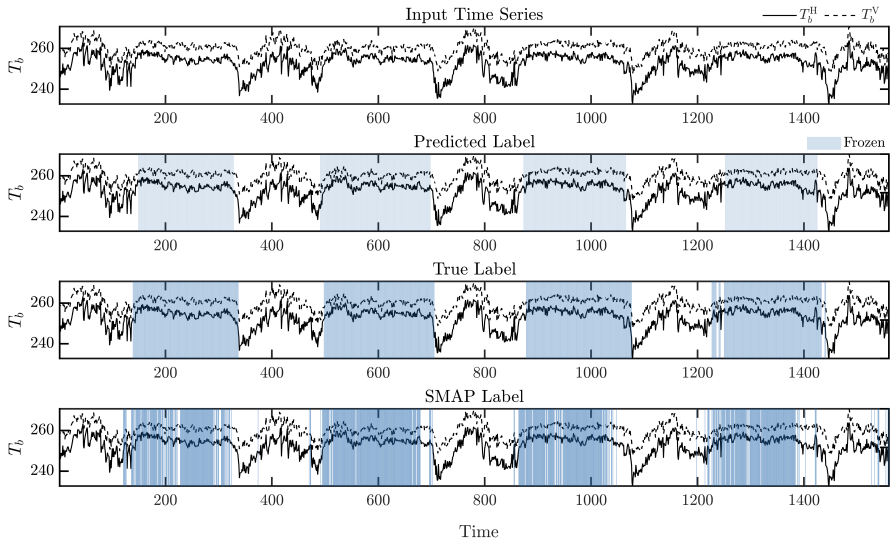
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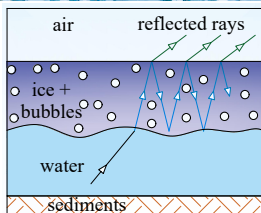
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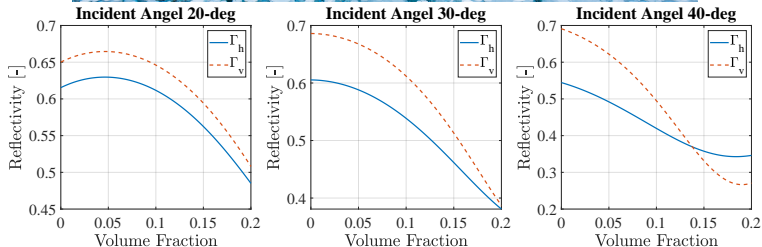


## Methane Ebullition Flux



$$\epsilon_m = \epsilon_i + \frac{3v_b\epsilon_i(\epsilon_b - \epsilon_i)}{(2\epsilon_i + \epsilon_b) - v_b(\epsilon_b - \epsilon_i)} \text{ (Tinga-Voss-Blossey)}$$

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## Summary

- ▶ Key ancillary variables that directly inform the RT model (e.g., clay fraction) can improve retrieval resolution.
- ▶ Not accounting for the effect of snow will lead to an overestimation of soil relative permittivity and VOD over the Arctic lands.
- ▶ Quantify NEE in wintertime when snow and ice cover the ground
- ▶ With  $10 \times 10$  km resolution, we can estimate the ice phenology of Arctic lakes greater than  $3 \times 3$  km.
- ▶ Methane ebullition flux to better understand the contribution of thawing permafrost on global warming.
- ▶ L-band PMW retrieval of active layer ice and carbon content?



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